



Comment on the paper „Field observations of soil moisture variability across scales” by Famiglietti et al.

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1. Abstract

15 In a recent paper, Famiglietti et al. (2008) analyzed more than 36,000 ground-based soil
moisture measurements to characterize soil moisture variability across spatial scales ranging
from 2.5 m to 50 km. They concluded that the relationship between soil moisture standard
deviation versus mean moisture content, $\sigma_{\theta}(\langle\theta\rangle)$, has a convex upward behaviour with
maximum values occurring at mean moisture contents of $0.17 \text{ cm}^3/\text{cm}^3$ and $0.19 \text{ cm}^3/\text{cm}^3$
20 for the 800 m and 50-km scale, respectively. Based on these data, they derived empirical
relationships between the coefficient of variation and the mean soil moisture content in order
to estimate the uncertainty in field observations of mean moisture content. The authors are to
be commended for providing this valuable database to the scientific community. We agree
with the authors that such data are important in improving our understanding about the
25 importance of sub-grid moisture variability in the parameterization and simulation of land
surface processes. However, the authors limited themselves to an empirical description of the
observed data by fitting exponential relationships to the mean moisture content versus CV
data. We feel that this is a missed opportunity and would like to argue that an interpretation
based on established theories and concepts in soil hydrology and upscaling theories could
30 provide alternative methods and new insights for interpreting such data sets.

2. Soil hydraulic properties and soil moisture variability

35 It can be shown from soil physical concepts that for a homogeneous soil, the shape of the
moisture retention curve can largely explain observed variations in surface soil moisture, at
any specific observation scale. For heterogeneous soils, stochastic upscaling theories may be
used to relate $\sigma_\theta(\langle\theta\rangle)$ to spatial variability in soil hydraulic properties. These theories can be
used to predict $\sigma_\theta(\langle\theta\rangle)$ and to examine the sensitivity of this function with respect to soil
40 hydraulic properties.

To better illustrate the potential contributions of the soil water retention curve on spatial
variations of surface soil moisture for homogeneous soil, we present the hysteretic soil water
retention curves across a range of soil textures in Fig. 1. The soil water retention curve defines
45 the unique relation between soil water potential, expressed by soil water pressure head (h, cm)
and soil water content (θ , $\text{cm}^3 \text{cm}^{-3}$), as determined by the soil's pore size distribution. The
functional relationships used to describe the curves were introduced by van Genuchten
(1980), with corresponding parameters listed in Table 1, representative for a sand, silt and
clay, as determined by Carsel and Parrish (1988). We chose to present the curves using a
50 logarithm scale, to better illustrate the soil's water retention in the dry range. As shown in
Fig.1, soil water content variations are expected to be the largest for intermediate values of
soil water content, θ , thus providing for a simple soil physical explanation for the upward
concave shape of Fig. 1 of Vereecken et al. (2007), and of Fig. 6 in Famiglietti et al. (2008).
Variability in h for a uniform soil may result from spatial variations in the soil moisture
55 regime, such as by plant water uptake, evaporation/infiltration and fluctuating water tables.

In this comment, we like to further impress the notion that observed spatial variations in field
soil water content can be partially explained by the shape of the soil water retention curve,
with the dependence partly determined by the slope of the retention curve, $d\theta/dh$, also known
60 as the soil water capacity, $C \text{ (cm}^{-1}\text{)}$. For the van Genuchten relationship, it is given by

$$\frac{d\theta}{dh} = (\theta_s - \theta_r) \left[1 + (\alpha|h|)^n \right]^{(1/n-2)} (\alpha|h|)^{n-1} \alpha(n-1) \quad , [1]$$

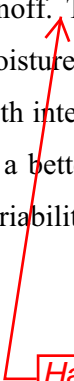
65 where θ_r and θ_s denote the residual and saturated soil water content, and α and n are curve
parameters (van Genuchten, 1980). In addition to presenting the soil water capacity in
(solid line), we show two additional curves that can partly explain the typical concave
associated with spatial variations of surface soil moisture. The second curve in Fig. 2,
by the dashed lines, shows $\Delta\theta/\Delta h$ as a function of θ , computed numerically, with Δh
equal to h , and centered around h . For example, at $h = -1000$ cm, the range of Δh is between -
500 and -1,500 cm, with a corresponding $\Delta\theta$ that is associated with these h -values. This
increase in Δh with h is typically observed in the field where the largest spatial variations in
soil water potential occur in the dry range, and it results in a shift of the curves to the left,
with maximum variation values depending on soil texture. To further illustrate the relevance
75 of the soil moisture retention curve, we added a third curve (dashed-dotted lines), to show the
additional effect of hysteresis of soil water retention (drying curve) on spatial water content
variation. These curves show the difference in h between the main drying and wetting curves,
 Δh -hysteresis, as a function of the mean soil water content, and illustrate that spatial
variations in soil water content are expected to show an upward concave curve, caused by
80 spatial variations of the wetting or drying regime of the surface soil. We hope that the
presented illustration makes a clear case that soil physical concepts can be used to explain

Conceptually, this situation would occur in a homogeneous soil with spatially variable infiltration.

Again, conceptually this situation may occur in a homogeneous soil with spatially variable infiltration.

variations in surface soil moisture across spatial scales even for a homogeneous soil.
of heterogeneous soils, it is well known that also the variability in the parameters
ture retention characteristic play an important role in determining soil moisture
variability (e.g. Vereecken et al., 2007a). Numerical simulations of soil moisture variability at
different degrees of saturation in heterogeneous unsaturated porous media were performed by
Roth (1995) and Harter and Zhang (1999) amongst others. Their results show that soil
moisture variability peaks at medium soil moisture content values. Closed form expressions
90 for the relationship between soil moisture variance and the statistical properties of soil
hydraulic parameters were derived by e.g. Russo et al. (1998) for steady state unsaturated
flow using the Gardner-Russo model of the moisture retention characteristic. An overview of
the state of the art in using stochastic methods for unsaturated flow in heterogeneous soils was
given by Zhang (2002). Recently, Vereecken et al. (2007) used results from stochastic
95 analysis of unsaturated flow in heterogeneous soils obtained by Zhang et al. (1998) to predict
the observed convex upward shapes of $\sigma_\theta(\langle\theta\rangle)$ also reported by Famiglietti et al. (2008).
Using this relationship for eleven textural classes, Vereecken et al. (2007b) showed that the
standard deviation of soil moisture peaked between 0.17 and 0.23 cm^3/cm^3 for most textural

classes. In addition, the parameter describing the pore-size distribution of soils controlled the
100 maximum value of the soil moisture standard deviation. The mean soil moisture values at
which the maximum soil moisture variability occurs are in very good agreement with the
values obtained by Famiglietti et al. (2008) from their very large database. This indicates the
potential value of stochastic theories of soil water processes in explaining and predicting the
observed spatial variability of soil moisture across scales. In this respect, we would like to
105 argue that $\sigma_{\theta}(\langle\theta\rangle)$ can be considered as a fundamental property of a heterogeneous soil,
which is related to the spatial variability in the moisture retention characteristic. Perturbations
of the $\sigma_{\theta}(\langle\theta\rangle)$ relationship may be caused by spatially and temporally heterogeneous fluxes
and sink/sources such as infiltration, evaporation, root water uptake, evaporation and surface
runoff. Taking stochastic theory as a starting point for the interpretation of observed soil
110 moisture variability and integrating and further developing upscaling approaches combined
with integrating knowledge from the fields of remote sensing and hydrology may finally lead
to a better understanding and a more fundamental interpretation of the role of soil moisture
variability in land surface processes across scales.



*Have you guys
looked at
Famiglietti's data to
see, whether we
can deduce the
stochastic
parameters from it?*

115 **References**

Carsel, R.F., and R.S. Parrish (1988). Developing joint probability distributions of soil water retention characteristics. *Water Resour. Res.* 24: 755-769.

120 Famiglietti, J. S., D. Ryu, A. A. Berg, M. Rodell, and T. J. Jackson (2008), Field observations of soil moisture variability across scales, *Water Resour. Res.*, 44, W01423, doi:10.1029/2006WR005804

Harter, T. and D. X. Zhang (1999). Water flow and solute spreading in heterogeneous soils with spatially variable water content. *Water Resour. Res.*, 35: 415-426.

125 Roth, K. (1995). Steady-State Flow in an Unsaturated, 2-Dimensional, Macroscopically Homogeneous, Miller-Similar Medium. *Water Resour. Res.*, 31: 2127-2140.

Russo, D. (1998). Stochastic analysis of flow and transport in unsaturated heterogeneous porous formation: effects of variability in water saturation. *Water Resour. Res.* 34: 569-581.

Scott, P. S., G. J. Farquhar, and N. Kouwen (1983). Hysteresis effects on net infiltration, *Advances in Infiltration, Publ. 11-83*, pp.163-170, Am. Soc. Agri. Eng., St. Joseph, Mich.

130 Van Genuchten, M. Th. (1980). A closed-form equation for pre-dicting the hydraulic conductivity of unsaturated soils. *Soil Sci. Soc. Am. J.* 44: 892-898.

135 Vereecken, H., R. Kasteel, J. Vanderborght and T. Harter (2007a). Upscaling hydraulic properties and soil water flow processes in heterogeneous soils: A review. *Vadose Zone Journal*, 6(1): 1-28.

Vereecken H., T. Kamaï, T. Harter, R. Kasteel, J.W. Hopmans, and J. Vanderborght, (2007b). Explaining soil moisture variability as a function of mean soil moisture: A stochastic unsaturated flow perspective. *Geophysical Research Letters* 34(22): Art. No. L22402

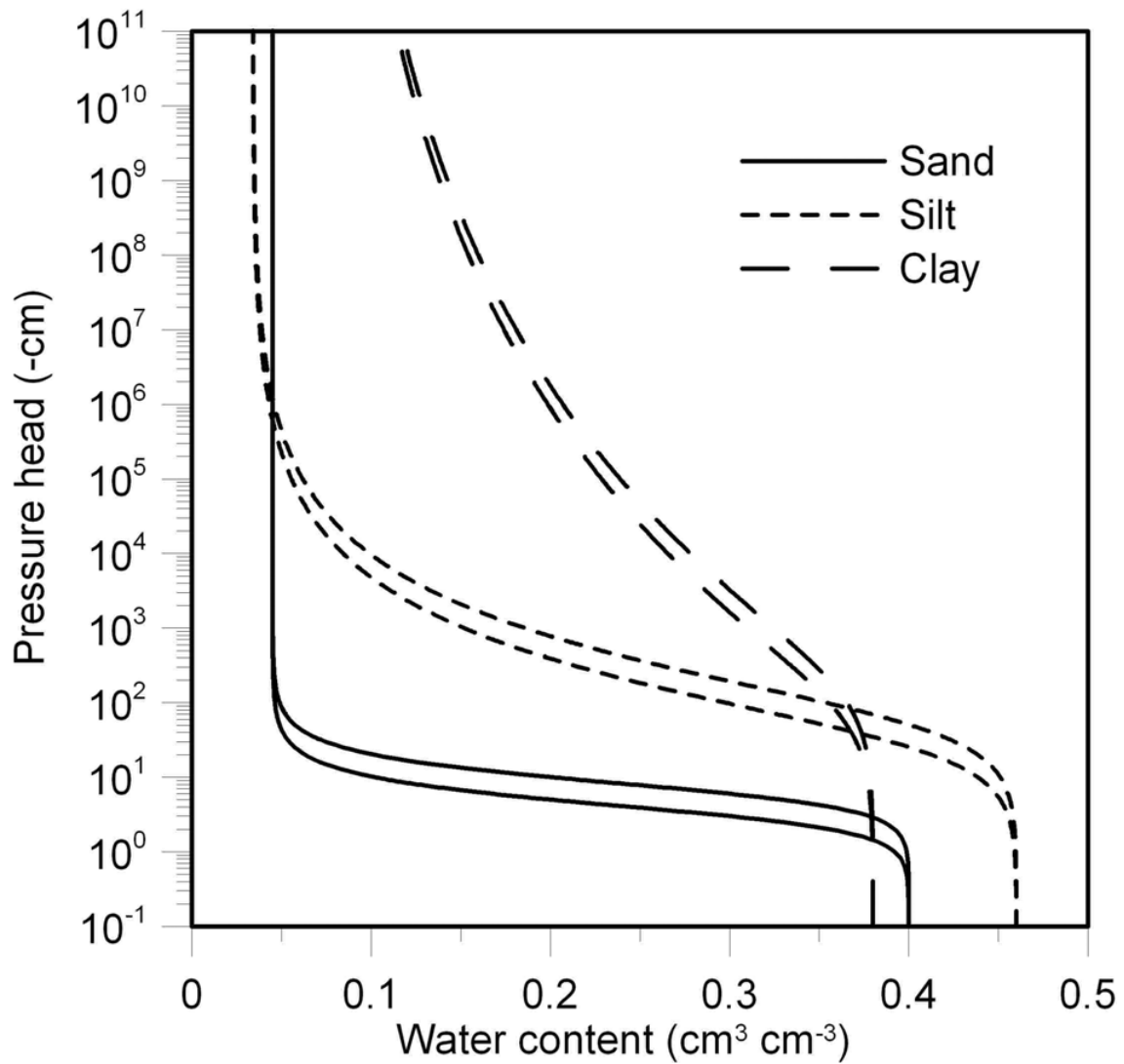
140 Yeh, T. C. J., L. W. Gelhar, and A.L. Gutjahr (1985). Stochastic-Analysis of Unsaturated Flow in Heterogeneous Soils .1. Statistically Isotropic Media. *Water Resour. Res.*, 21: 447-456.

Zhang, D. X., T. C. Wallstrom, and C.L. Winter (1998). Stochastic analysis of steady-state unsaturated flow in heterogeneous media: Comparison of the Brooks-Corey and Gardner-
145 Russo models. *Water Resour. Res.*, 34: 1437-1449.

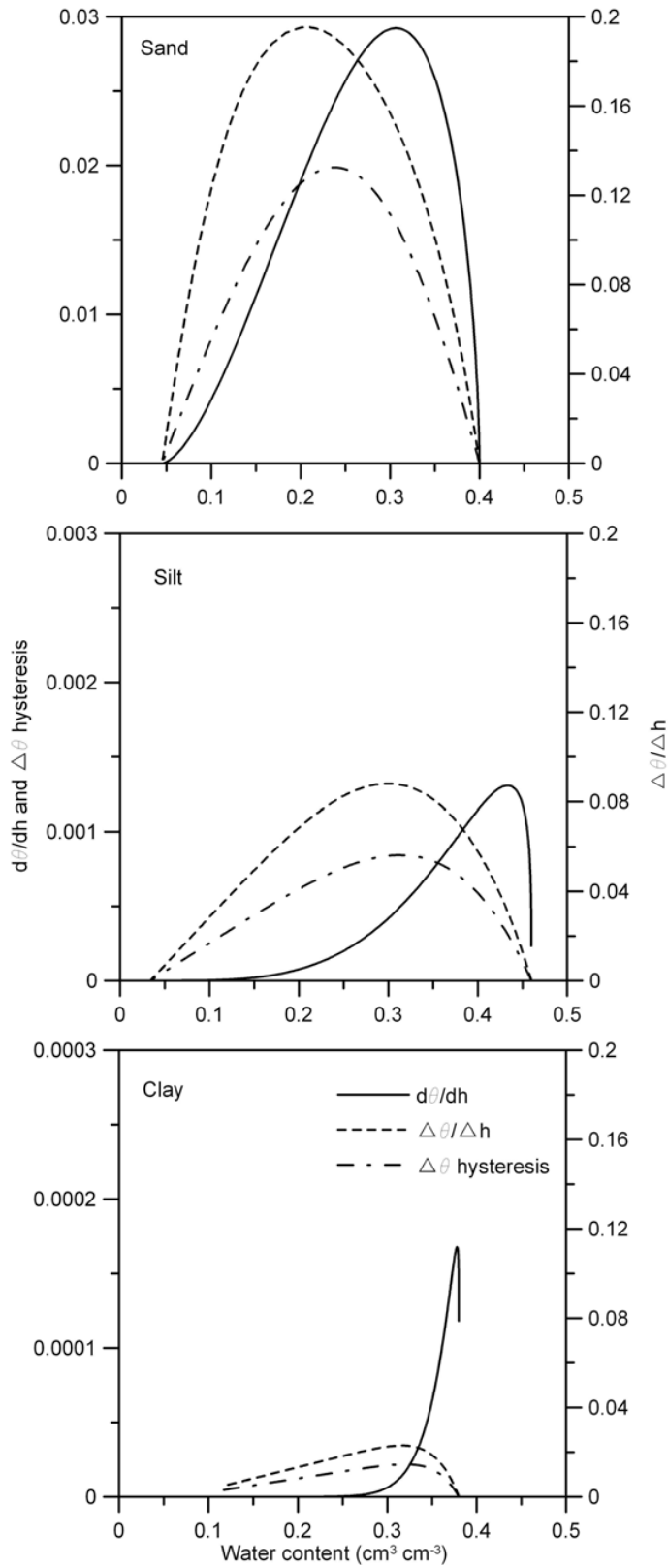
Zhang, D.X. (2002). *Stochastic methods for flow in porous media: coping with uncertainties*. Academic Press, p. 368.

Table 1. Van Genuchten parameters (Carsel and Parrish (1988))

	θ_r	θ_s	$\alpha_{\text{ dry (cm}^{-1}\text{)}}$	$\alpha_{\text{ wet (cm}^{-1}\text{)}}$	n
Sand	0.045	0.4	0.145	0.29	2.68
Silt	0.034	0.46	0.016	0.032	1.37
Clay	0.068	0.38	0.008	0.016	1.09



150 Figure 1. Retention curves with hysteresis for the three soils in Table 1. Main drying and wetting curves are represented by top and bottom curves, respectively, for each soil.



155 Figure 2. The slope of the retention curve, $d\theta/dh$, $\Delta\theta/\Delta h$, and Δh -hysteresis, as a function of mean soil water content, θ . Note the different scales between the left and right vertical axis.

